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# Probabilistic physics-informed neural network

## for seismic petrophysical inversion

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## ABSTRACT

The main challenge in the inversion of seismic data to predict the petrophysical properties of hydrocarbon-saturated rocks is that the physical relations that link the data to the model properties are often non-linear and the solution of the inverse problem is generally not unique. As a possible alternative to traditional stochastic optimization methods, we propose to adopt machine learning algorithms by estimating relations between data and unknown variables from a training dataset with limited computational cost. We present a probabilistic approach for seismic petrophysical inversion based on physics-informed neural network with a reparameterization network. The





distributions of the model variables and data errors. Monte Carlo methods (Mosegaard and Tarantola, 1995; Bosch et al., 2007; Connolly and Hughes, 2016; Sen and Biswas, 2017; Zhu and Gibson, 2018; De Figueiredo et al., 2019a,b; Stuart et al., 2019) allow estimating the unknown posterior distribution of the model variables by sampling from a proposal distribution. Stochastic optimization algorithms such as genetical algorithms, simulating annealing and particle swarm optimizations (Sen and Stoffa, 2013; Aster et al., 2018; Grana et al., 2021) can also be used to estimate the solution of the inverse problem. Monte Carlo methods and stochastic optimization algorithms are generally applied to non-linear inverse problems and/or when the distribution of the model variables is not Gaussian; however, these methods are more computationally expensive than analytical methods. Furthermore, the uncertainty assessment might not be precise if a prior distribution of the model variables is not chosen accurately.

In recent years, machine learning (ML) or deep learning (DL) methods have emerged as a possible alternative (Dramsch, 2020; Bhattacharya, 2021; Bhattacharya and Di, 2022) for automatically mapping relations between measured data and model variables without significant computational cost and prior assumptions. In particular, supervised learning methods allow us to derive the relation between the data and the properties of interest from training data without the

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3 use of a physical forward model. Several deep learning algorithms have been adopted in  
4 exploration geoscience applications, including fault detection (Hale, 2013; Xiong et al., 2018; Wu  
5 et al., 2018), facies classification (Matos et al., 2007; Wrona et al., 2018; Alaudah et al., 2019; Liu  
6 et al., 2020), porous media modeling (Kamrava et al., 2020, 2021), seismic inversion (Das et al.,  
7 2018; Mosser et al., 2018; Alfarraj and AlRegib, 2019a,b), and petrophysical characterization  
8 (Shahraeeni and Curtis, 2011; Shahraeeni et al., 2012; Das and Mukerji, 2020; Verma et al., 2021;  
9 Bhattacharya, 2022; Vashisth and Mukerji, 2022). Furthermore, deep learning methods have been  
10 combined with probability theory to quantify the model uncertainty (Sengupta et al., 2020; Feng  
11 et al., 2021; Yang et al., 2021; Daw et al., 2021; Mosser and Naeini, 2022).

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13 The application of ML and DL algorithms to geophysical applications still has several  
14 challenges such as insufficient training data to fit the complex relations between observations and  
15 target model variables. For example, in seismic petrophysical inversion, the size of the seismic  
16 data (observations) is orders of magnitude larger than the dimension of labeled model variables  
17 from well logs measurements (target model variables) due to the high cost of drilling wells,  
18 especially in the exploration area. Another challenge is that the training dataset might be biased by  
19 the location of the wells which might not account for all possible rock and fluid types. In addition,  
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 3 many deep learning methods do not account for uncertainty quantification due to the high accuracy  
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 5 of the most likely models, however, in geophysical inverse problems data errors and model  
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 7 heterogeneity generally lead to high uncertainty in model predictions.  
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 10 We propose a probabilistic physics-informed neural network (P-PINN) algorithm for seismic  
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 12 petrophysical inversion with the goal of estimating the most likely model of petrophysical  
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 14 properties, the unknown hyperparameters of the rock physics model, and the uncertainty of the  
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 16 model predictions. Physics-informed neural networks (PINN) method is a subclass of deep  
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 18 learning algorithms that integrate data and physical mathematical models and implement them  
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 20 through neural networks or other kernel-based regression networks (Karniadakis et al., 2021).  
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 22 Unlike the work by Karniadakis et al. (2021), the proposed approach does not adopt partial  
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 24 differential equations, since the physics is based on geophysical models consisting of algebraic  
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 26 equations. Compared to other ML or DL algorithms, PINN algorithms have the advantage of  
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 28 incorporating prior knowledge and/or physical constraints into the networks. PINN algorithms  
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 30 have already been used in geophysical inversion for acoustic impedance inversion (Alfarraj and  
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 32 AlRegib, 2019a) and elastic impedance inversion (Alfarraj and AlRegib, 2019b; Biswas et al.,  
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 34 2019). Furthermore, PINN have been used for solving partial differential equations (PDE) in  
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stack seismic data) and multivariate model setting (e.g., prediction of petrophysical properties from pre-stack seismic data).

## THEORY

The general form of the inverse problem aims at predicting the model variables  $\mathbf{m}$  from measured data  $\mathbf{d}$  according to the governing physical operator  $\mathcal{F}$ :

$$\mathbf{d} = \mathcal{F}(\mathbf{m}) + \mathbf{e}, \quad (1)$$

where  $\mathbf{e}$  represents the measurement errors. In non-linear optimization methods, the goal is to find the value  $\hat{\mathbf{m}}$  that minimizes the misfit between the measured data  $\mathbf{d}$  and geophysical model prediction  $\hat{\mathbf{d}}$ :

$$\hat{\mathbf{m}} = \operatorname{argmin}_{\mathbf{m}} (\|\mathbf{d} - \hat{\mathbf{d}}\|). \quad (2)$$

In the context of inverse problems, deep learning methods are used to find an approximation of the inverse operator  $\mathcal{G} \cong \mathcal{F}^{-1}$  and minimize the difference between the predicted output  $\hat{\mathbf{m}}$  and the target variables  $\mathbf{m}$  on a training dataset:

$$\hat{\mathbf{m}} = \mathcal{G}(\mathbf{d}; \boldsymbol{\theta}) \quad (3)$$

where  $\boldsymbol{\theta}$  represents the training parameters in the neural networks (i.e., weights and biases).

We propose a new method based on a probabilistic approach to PINN, namely P-PINN for seismic petrophysical inversion that combines the prediction of the most likely model and the estimation of the rock physics model hyperparameters as well as their uncertainty. The architecture includes a physics-informed deep neural network that consists of an inverse network and a forward model, and a reparameterization network (Figure 1). In this setting,  $\mathbf{d}$  denotes the seismic data,  $\mathbf{m}_L$  denotes the petrophysical properties from well-log data (labeled data),  $\mathbf{d}_L$  denotes the corresponding seismic data with respect to  $\mathbf{m}_L$ ,  $\boldsymbol{\lambda}$  denotes rock physics model hyperparameters, and  $\hat{\mathbf{d}}$ ,  $\hat{\mathbf{m}}_L$ ,  $\hat{\boldsymbol{\lambda}}$  denote the estimated seismic data, petrophysical properties and rock physics model hyperparameters, respectively.

### Physics-informed neural network

To find the optimal estimation of  $\mathcal{G} \cong \mathcal{F}^{-1}$ , the training parameters  $\boldsymbol{\theta}$  are learned either in a supervised or unsupervised manner. In the context of seismic petrophysical inversion, supervised

learning requires a training dataset consisting of a set of seismic and petrophysical property pairs [ $\mathbf{d}_L, \mathbf{m}_L$ ]. The learning process is performed by minimizing the petrophysical loss function  $L_p(\boldsymbol{\theta})$

between the predicted model variables  $\hat{\mathbf{m}}_L$  and target (labeled) model variables  $\mathbf{m}_L$ :

$$L_p(\boldsymbol{\theta}) = \sum_{i=1}^n C(\mathcal{G}(\mathbf{d}_L; \boldsymbol{\theta}), \mathbf{m}_L) \quad (4)$$

where  $n$  represents the number of training samples and  $C$  represents the dissimilarity metric, such as the mean squared error (MSE) in a regression problem or cross-entropy in a classification problem.

In the proposed PINN, the physical constraints are embedded in the neural network by adding a seismic loss function  $L_s(\boldsymbol{\theta})$  based on the forward model  $\mathcal{F}$ . Seismic data are the input of the neural network to obtain estimates of the model variables (i.e.,  $\mathcal{G}(\mathbf{d}; \boldsymbol{\theta})$ ), that are then the input of the forward model  $\mathcal{F}$  to obtain the predictions of the seismic data (i.e.,  $\mathcal{F}(\mathcal{G}(\mathbf{d}; \boldsymbol{\theta}))$ ). The misfit between the measured and predicted seismic data is then the additional loss term  $L_s(\boldsymbol{\theta})$ :

$$L_s(\boldsymbol{\theta}) = \sum_{i=1}^n C(\mathcal{F}(\mathcal{G}(\mathbf{d}; \boldsymbol{\theta})), \mathbf{d}). \quad (5)$$

The total loss function can then be written as:

$$L_t(\boldsymbol{\theta}) = \alpha L_p(\boldsymbol{\theta}) + \beta L_s(\boldsymbol{\theta}) \quad (6)$$

where  $\alpha$  and  $\beta$  are tuning parameters to control the influence of each loss function. The neural network is trained by minimizing the total loss function  $L_t(\boldsymbol{\theta})$ . After sufficient training, an optimal function  $\mathcal{G}$  is built to map measurements  $\mathbf{d}$  into model variables  $\mathbf{m}$ . The operator  $\mathcal{G}$  is used for the model variables predictions  $\mathbf{m}_p$  when the unseen measurements  $\mathbf{d}_p$  are given to neural networks.

$$\mathbf{m}_p = \mathcal{G}(\mathbf{d}_p; \boldsymbol{\theta}). \quad (7)$$

### Model hyperparameter estimation

One of the advantages of PINN is that it can provide not only an estimate of the optimal values of the model variables but also an estimate of rock physics model hyperparameters of the forward model. For example, some rock physics parameters  $\boldsymbol{\lambda}$  in the forward operator, such as critical porosity and coordination number in granular media models or aspect ratio in inclusion models (Mavko et al., 2020), might be unknown in certain applications. These parameters are often defined

based on indirect measurements, core sample images, or nearby field studies. In this work, we estimate the rock physics model hyperparameters from the data simultaneously with the inverse mapping approximation  $\mathcal{G}(\mathbf{d};\boldsymbol{\theta},\boldsymbol{\lambda})$  within the PINN. Hence, the loss function in Equation (5) is reformulated as:

$$L_s(\boldsymbol{\theta},\boldsymbol{\lambda}) = \sum_i^n L(\mathcal{F}(\mathcal{G}(\mathbf{d};\boldsymbol{\theta},\boldsymbol{\lambda})),\mathbf{d}), \quad (8)$$

and the total loss function in Equation (6) can be rewritten as:

$$L_t(\boldsymbol{\theta},\boldsymbol{\lambda}) = \alpha L_p(\boldsymbol{\theta}) + \beta L_s(\boldsymbol{\theta},\boldsymbol{\lambda}). \quad (9)$$

Then the total loss  $L_t(\boldsymbol{\theta},\boldsymbol{\lambda})$  is minimized by training the neural network and the model variables predictions are obtained by applying the operator  $\mathcal{G}$  with the optimal parameters  $\boldsymbol{\lambda}$ :

$$\mathbf{m}_p = \mathcal{G}(\mathbf{d}_p;\boldsymbol{\theta},\boldsymbol{\lambda}). \quad (10)$$

### Bayesian formulation

To quantify the uncertainty associated with the neural network, a probabilistic approach can

be applied by formulating the estimation of the parameters  $\theta$  of the neural network  $P(\theta | \mathbf{D})$  given the data  $\mathbf{D}$  in a Bayesian framework as :

$$P(\theta | \mathbf{D}) = \frac{P(\mathbf{D} | \theta)P(\theta)}{P(\mathbf{D})}. \quad (11)$$

When the forward model includes unknown hyperparameters  $\lambda$ , equation 11 can be rewritten to compute the posterior probability distribution  $P(\theta, \lambda | \mathbf{D})$  of the neural parameters  $\theta$  and hyperparameters  $\lambda$ :

$$P(\theta, \lambda | \mathbf{D}) = \frac{P(\mathbf{D} | \theta, \lambda)P(\theta, \lambda)}{P(\mathbf{D})} = \frac{P(\mathbf{D} | \theta, \lambda)P(\theta)P(\lambda)}{P(\mathbf{D})} \quad (12)$$

where we assume that the neural parameters  $\theta$  and hyperparameters  $\lambda$  are independent.

We address the estimation of the posterior distribution in equation 12 using the Approximate Bayesian Computation (ABC). ABC is a class of Bayesian statistics methods which approximate the likelihood function with Monte Carlo simulation, and the outcomes of estimations are accepted or rejected according a rejection algorithm (Beaumont et al., 2002; Csilléry et al., 2010; Sunnåker et al., 2013; Fernández et al., 2022). In the proposed approach, at each simulation, the parameters  $\theta$  and  $\lambda$  are first sampled from the prior distribution, a prediction  $\hat{\mathbf{D}}$  given  $\theta$  and  $\lambda$  is then obtained

through the neural network model.  $\hat{\mathbf{D}}$  is accepted if the acceptance criterion is satisfied:

$$\rho_d(\hat{\mathbf{D}}, \mathbf{D}) \leq \epsilon \quad (13)$$

where  $\rho_d$  measures the distance between  $\hat{\mathbf{D}}$  and  $\mathbf{D}$  based on a given metric and the threshold  $\epsilon$  represents a tolerance value. The posterior uncertainty is generally affected by the tolerance value  $\epsilon$ . By running MC simulations, the posterior distribution is approximated from the accepted models. ABC can be also used for Bayesian model comparison by calculating the posterior ratio of the models and defining a parameter  $K$  that assesses the plausibility of two different models  $M_1$  and  $M_2$ :

$$K = \frac{P(M_1 | \mathbf{D})}{P(M_2 | \mathbf{D})} \quad (14)$$

The posterior ratio  $K$  is related to the Bayes factor if the prior models of  $P(M_1)$  and  $P(M_2)$  are equal.

The ABC method is based on a Monte Carlo simulation that does not require the explicit evaluation of the likelihood function, which is a great advantage especially for the cases where the likelihood model is analytically intractable. The main limitation of the proposed approach is that



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3 the posterior uncertainty might depend on the tolerance value. If the tolerance value is too large,  
4 the method tends to accept a large number of models (i.e., the prior model is dominant with respect  
5 to the data), whereas if the tolerance value is too small, the method tends to accept a small number  
6 of models (i.e., the data are dominant with respect to the prior model).  
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### 10 **Network architecture**

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12 In the context of geophysical inverse problems, the inverse problem aims to estimate the  
13 relation between the measured data and the petrophysical model variables, and the forward model  
14 aims to embed the seismic and rock physics equation in an unsupervised loss function. The inverse  
15 model consists of two deep neural networks: an inverse network (Figure 2a) and a  
16 reparameterization network (Figure 2b), for the estimation of petrophysical model variables and  
17 of the rock physics model hyperparameters respectively. If the model parameters can be accurately  
18 assessed from core measurements, the reparameterization network can be simplified or removed.  
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27 The inverse network is extended from the network proposed in Alfarraj and AlRegib (2019a,b)  
28 (Figure 2a). The inverse network includes four blocks: Sequential block, Local pattern block,  
29 Upscaling block, and Regression block. The Sequential block is designed with three layers of gated  
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3 recurrent units (GRUs) to capture the low-frequency trend in the input-output relations. The Local  
4 pattern block is designed with three 1D convolutional layers followed by a 1D convolutional layer  
5 to capture the high-frequency content in the training data. The Upscaling block is designed with  
6 two deconvolution layers to balance the resolution mismatch between seismic data (unlabeled) and  
7 well (labeled) data. The Regression block is designed with a GRU followed by a fully connected  
8 (FC) layer to regress the generated feature from previous layers to the target domain.  
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13 In addition to the inverse network, a reparameterization network is also included in the inverse  
14 model. The reparameterization network consists of three FC layers (Figure 2b). In the case of  
15 seismic petrophysical inversion, the rock physics model hyperparameters are estimated through  
16 this neural network and the outputs are passed to the forward model.  
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## APPLICATION

To validate the proposed method, we test the inversion on a dataset from an oil-saturated clastic  
reservoir in the North Sea. The dataset includes well-log data and 2D simulated maps of porosity,  
clay volume, and water saturation (Dvorkin et al., 2014). The dataset represents a clastic reservoir  
with a sequence of oil-saturated sandstone layers alternating with non-reservoir shale layers. A  $0^\circ$



are shown in Figure 4 and include 337 traces. We randomly extract a subset of 30 traces from both seismic data and porosity model as the input-target pairs of the training dataset and split them into two groups of 24 traces for training and 6 traces for validation, respectively. The target petrophysical properties are converted into the time domain for consistency with seismic data.

We first demonstrate the application of the P-PINN network and compare it to a traditional neural network (NN) without physical constraints. The same training parameters are used for the two training processes with the only difference that a physics-informed loss function is added to the P-PINN network. In the ABC, we adopt the Mean Absolute Error (MAE) between true properties and predicted properties as the threshold for the acceptance / rejection criterion. We then run 500 simulations and the models that satisfy the threshold ( $\rho_{dpor} < 0.015$ ) are accepted and the others rejected. In this application, 92 models are accepted with the regular NN and 149 models are accepted using the P-PINN with Bayes factor  $K = 1.62$ , showing the higher relative plausibility of the P-PINN model. Figure 5 illustrates the comparison of the 1D predictions of the regular NN (Figure 5a) and P-PINN (Figure 5b). The predicted results obtained with the P-PINN show an improvement in terms of prediction accuracy and a smaller uncertainty range. We calculated three metrics, correlation coefficient, the determination coefficient  $R^2$ , and the MAE, of

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in Figure 10. For simplicity, only the weights of the last regression layer are shown.

### Synthetic example: multivariate case

In the second example, we test the proposed method in a multivariate setting, where the model variables of interest are porosity, clay volume, and water saturation. The reference well-log data are shown in Figure 11. We adopt three incident angles ( $15^\circ$ ,  $30^\circ$ ,  $45^\circ$ ) to mimic the near, mid, and far angle stacks of the partially stacked seismic dataset. The elastic properties are calculated using the soft sand model with Gassmann's equation and the seismic data are computed by convolving the known source wavelet with the amplitude variation-versus-offset (AVO) approximation of the reflection coefficients (Aki and Richards, 2002). The wavelet is a Ricker wavelet with dominant frequency of 30 Hz. The seismic dataset with random noise is shown in Figure 12, with an average signal-to-noise ratio of 5. The true porosity section and synthetic seismic section are shown in Figure 13 and Figure 14. As in the previous example, 30 traces of input-target properties are extracted as the training dataset with a train-validation ratio of 0.8. In the ABC, we apply the following MAE threshold:  $\rho_{dpor} < 0.019$  and  $\rho_{dclay} < 0.095$  and  $\rho_{dsat} < 0.081$ , for applying acceptance / rejection. We run 100 simulations and 39 simulations are accepted using regular NN

and 69 simulations are accepted using P-PINN with Bayes factor  $K$  of 1.77.

Figure 15 illustrates the comparison of 1D predictions of the regular NN (Figure 15a) and P-PINN (Figure 15b) by applying ABC. Overall, the P-PINN provides more accurate results with a smaller difference between true and predicted properties. In particular, the prediction of water saturation is more accurate in the hydrocarbon layers. The comparison of the metrics distributions (correlation and determination coefficients and MAE) between true and accepted models of the three variables are shown in Figure 16 for both regular NN and from P-PINN. Overall, the most likely models from P-PINN have higher correlation coefficient, higher  $R^2$  and lower MAE compared to models obtained with the regular NN. Figures 17 and 18 illustrate the mean and standard deviation of the model variables obtained with NN and P-PINN, respectively. Overall, the P-PINN allows the prediction of both low and high frequency information from the data. The P-PINN results show a smaller standard deviation for all three variables, which indicates lower uncertainty in the prediction results.

In this example, we assume two unknown rock physics hyperparameters, namely the critical porosity and the coordination number and we estimate them using the P-PINN. To generate the reference synthetic seismic data, we used a critical porosity of 0.4 and a coordination number of

7. The estimated distributions of the two rock physics model hyperparameters are shown in Figure 19. The predicted coordination number is 6.92 with a standard deviation of 0.129 and the predicted critical porosity is 0.408 with a standard deviation of 0.0051. The posterior distributions of the network weights are shown in Figure 20 for regular NN and P-PINN. Only the weights from the last regression layer are demonstrated for simplicity.

### Additional examples

In this section, we present two additional applications. In the first example, we adopt the model in Figure 13 where we impose the oil-water contact (OWC) at 1.85 s to mimic the presence of oil- and water-saturated rocks within the main reservoir layer. A synthetic seismic dataset is generated with the same forward model used in the previous example (Figure 21). We apply the ABC with the following MAE thresholds:  $\rho_{d_{por}} < 0.02$  and  $\rho_{d_{clay}} < 0.11$  and  $\rho_{d_{sat}} < 0.05$ . We run 100 simulations: 31 simulations are accepted using regular NN and 36 simulations using P-PINN with Bayes factor  $K$  of 1.16. The comparison between the inverted properties from regular neural networks and from P-PINN is shown in Figures 22 and 23. Overall, both inversion results are accurate, however the P-PINN method provides lower standard deviations, which means less





We quantify the performance of the neural network methods using three metrics: correlation coefficient,  $R^2$ , and mean absolute error (MAE). The values of these metrics for the univariate case are shown in Table 1 and the values for the multivariate case are in Table 2. Compared to the regular NN, overall improvements for all the model variables obtained by the P-PINN are observed, i.e., higher correlation coefficient and  $R^2$  and lower MAE.

The proposed method is scalable and largely adaptable to different inverse problems by combining the main components of the model, namely the inverse network, reparameterization network, and forward model. For example, the proposed inverse network could be replaced by any other neural network such as convolutional neural network (CNN), recurrent neural network (RNN), Fully connected layers (FC layers), or a combination of them. Similarly, the reparameterization network could be replaced by constant values or spatial functions if the parameters are known or can be estimated from other data and the forward model can be substituted with any physical-mathematical relation. The method can be naturally extended to 3D and time-lapse datasets, but the uncertainty quantification study might be limited by the computational costs of multiple simulations for large datasets.

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## CONCLUSIONS

We presented a probabilistic approach to physics-informed neural network for seismic petrophysical inversion for the simultaneous estimation of the unknown model variables and the rock physics model hyperparameters. The probabilistic approach adopted for the PINN allows us to quantify the uncertainty in the model variables predictions and parameters estimation. The uncertainty quantification approach is formulated in a Bayesian framework where the posterior distribution of the training parameters (i.e., weights and biases) and the forward model parameters (e.g., rock physics parameters) are assumed to be probability distributions. The posterior distribution is computed using Bayes' rule by sampling the training and forward model parameters from the prior and accepting or rejecting the models using the ABC approach. The novelty of the method is the implementation of an additional neural network for hyperparameter estimation in the PINN network and ABC for uncertainty quantification. The proposed method can be embedded in the network known or partially known physics relations and it is easily scalable. We validated the proposed method with univariate and multivariate models and in both cases, the P-PINN provide more accurate and more precise (i.e., less uncertain) predictions. In both cases, the rock physics model hyperparameters are accurately estimated. The presented examples show the



$$\mu_{HM} = \frac{5 - 4\nu_{sol}}{5(2 - \nu_{sol})} \sqrt[3]{\frac{3c^2(1 - \phi)^2 \mu_{sol}^2}{2\pi^2(1 - \nu_{sol})^2}} P$$

$$K_{dry} = \left( \frac{\frac{\phi}{\phi_c}}{K_{HM} + \frac{4}{3}\mu_{HM}} + \frac{1 - \frac{\phi}{\phi_c}}{K_{sol} + \frac{4}{3}\mu_{HM}} \right)^{-1} - \frac{4}{3}\mu_{HM}$$

$$\mu_{dry} = \left( \frac{\frac{\phi}{\phi_c}}{\mu_{HM} + \frac{\mu_{HM}}{6}\xi_{HM}} + \frac{1 - \frac{\phi}{\phi_c}}{\mu_{sol} + \frac{\mu_{HM}}{6}\xi_{HM}} \right)^{-1} - \frac{\mu_{HM}}{6}\xi_{HM}$$

$$\xi_{HM} = \frac{9K_{sol} + 8\mu_{sol}}{K_{sol} + 2\mu_{sol}}$$

(A1)

where  $K_{HM}$  and  $\mu_{HM}$  are the Hertz-Mindlin moduli and shear moduli at the critical porosity,  $\phi$  is the porosity,  $c$  represents coordinate number,  $\phi_c$  the critical porosity,  $\mu_{sol}$  the shear moduli of the solid phase,  $\nu_{sol}$  the Poisson's ratio of solid phase, and  $P$  is the effective pressure.

The model combines the dry moduli in equation A1 with Gassmann's equations to compute the bulk and shear moduli of the saturated rock  $K_{sat}$  and  $\mu_{sat}$ :

$$K_{sat} = K_{dry} + \frac{\left(1 - \frac{K_{dry}}{K_{sol}}\right)^2}{\frac{\phi}{K_{fl}} + \frac{(1 - \phi)}{K_{sol}} - \frac{K_{dry}}{K_{sol}^2}}$$

$$\mu_{sat} = \mu_{dry}$$

(A2)

where  $K_{fl}$  is the bulk modulus of the fluid phase. The P- and S- wave velocity  $V_p$  and  $V_s$  are then computed as:

$$V_p = \sqrt{\frac{K_{sat} + \frac{4}{3}\mu_{sat}}{\rho}}$$

$$V_s = \sqrt{\frac{\mu_{sat}}{\rho}}$$

(A3)

where  $\rho = (1 - \phi)\rho_{sol} + \phi\rho_{fl}$  is the density of the saturated rock and is computed using a mass balance of the density of the solid and fluid phases  $\rho_{sol}$  and  $\rho_{fl}$  weighted by porosity.

We then use the Aki Richards AVO equation to calculate the reflectivity coefficients at each interface time value as a function of the incident angle  $\theta$ :

$$r_{pp}(\theta) = \frac{1}{2}(1 + \tan^2 \theta) \frac{\Delta V_p}{V_p} - 4 \frac{\overline{V_s}^2}{V_p^2} \sin^2 \theta \frac{\Delta V_s}{V_s} + \frac{1}{2} \left(1 - 4 \frac{\overline{V_s}^2}{V_p^2} \sin^2 \theta\right) \frac{\Delta \rho}{\bar{\rho}}$$

(A4)

where  $\overline{V_p}$  is the average P-wave velocity across the interface,  $\overline{V_s}$  is the average S-wave velocity across the interface,  $\overline{\rho}$  is the average density across the interface,  $\Delta V_p$  is the P-wave velocity differential across the interface,  $\Delta V_s$  is the S-wave velocity differential across the interface, and  $\Delta \rho$  density differential across the interface.

Finally, the seismic signal  $s(\theta)$  for each incident angle  $\theta$  is computed as a convolution between the wavelet  $w(\theta)$  and reflectivity coefficients  $r_{pp}(\theta)$ :

$$s(\theta) = w(\theta) * r_{pp}(\theta).$$

(A5)

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		Correlation coefficient	R <sup>2</sup>	MAE
Porosity	NN	0.8612	0.7100	0.01411
	P-PINN	0.8717	0.7330	0.01404

Table 1: Comparison of performance metrics between the mean of the predictions from regular NN and from P-PINN of the univariate inverse problem.

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		Correlation coefficient	R <sup>2</sup>	MAE
Porosity	NN	0.777	0.563	0.018
	P-PINN	0.831	0.666	0.017
Clay volume	NN	0.712	0.432	0.082
	P-PINN	0.755	0.548	0.075
Water saturation	NN	0.840	0.668	0.069
	P-PINN	0.886	0.769	0.062

Table 2: Comparison of performance metrics between the mean of the predictions from regular NN and from P-PINN of the multivariate inverse problem.

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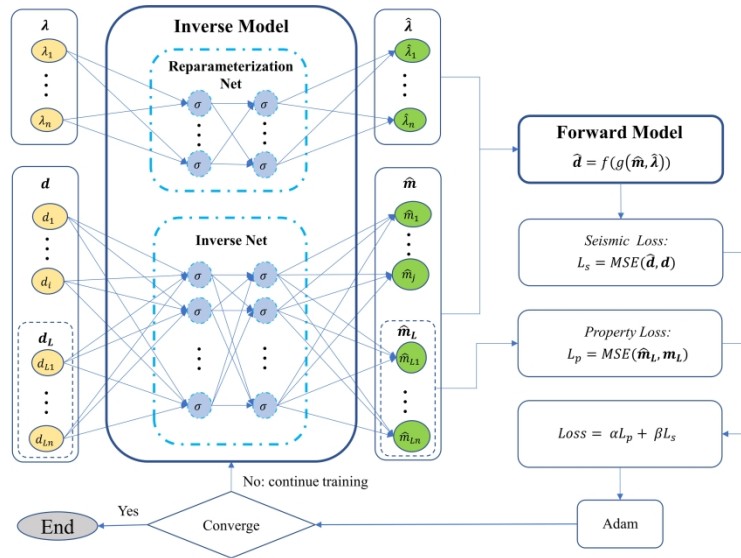


Figure 1: Workflow of the proposed method.

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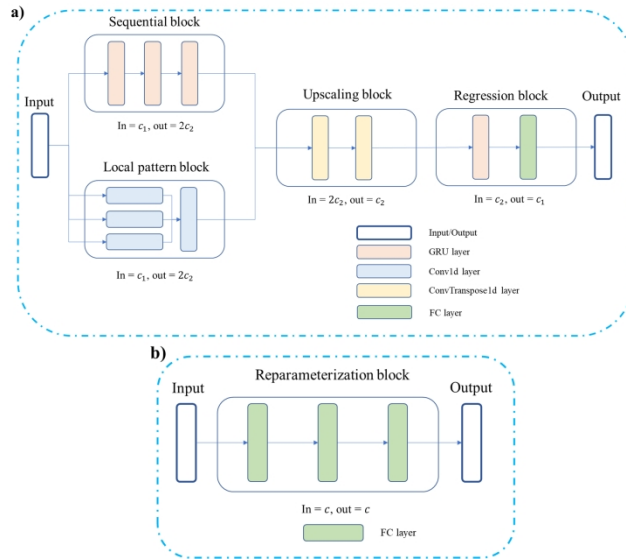


Figure 2: Neural network architectures, a): Inverse net, and b) Reparameterization net.

338x190mm (300 x 300 DPI)

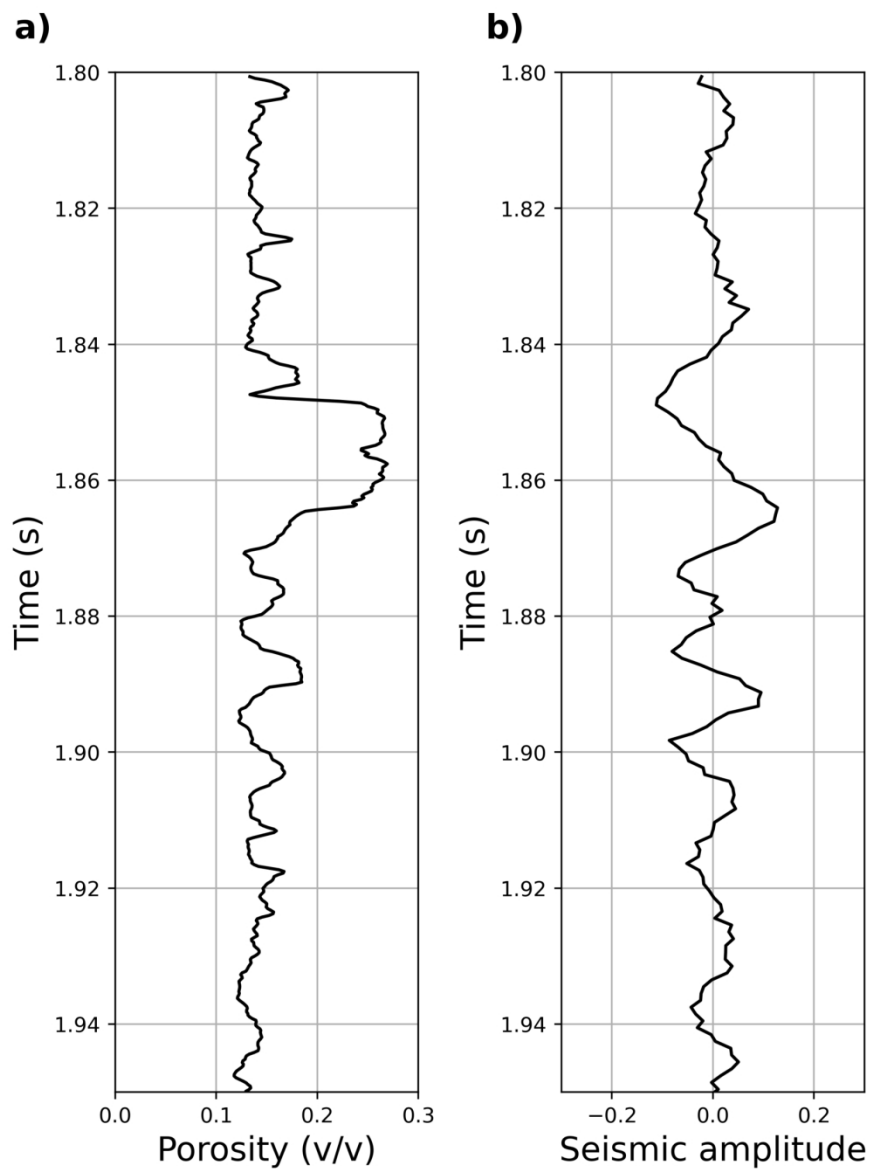


Figure 3: 1D dataset for the univariate inverse problem, a) porosity log, and b) synthetic seismic data.

152x203mm (300 x 300 DPI)

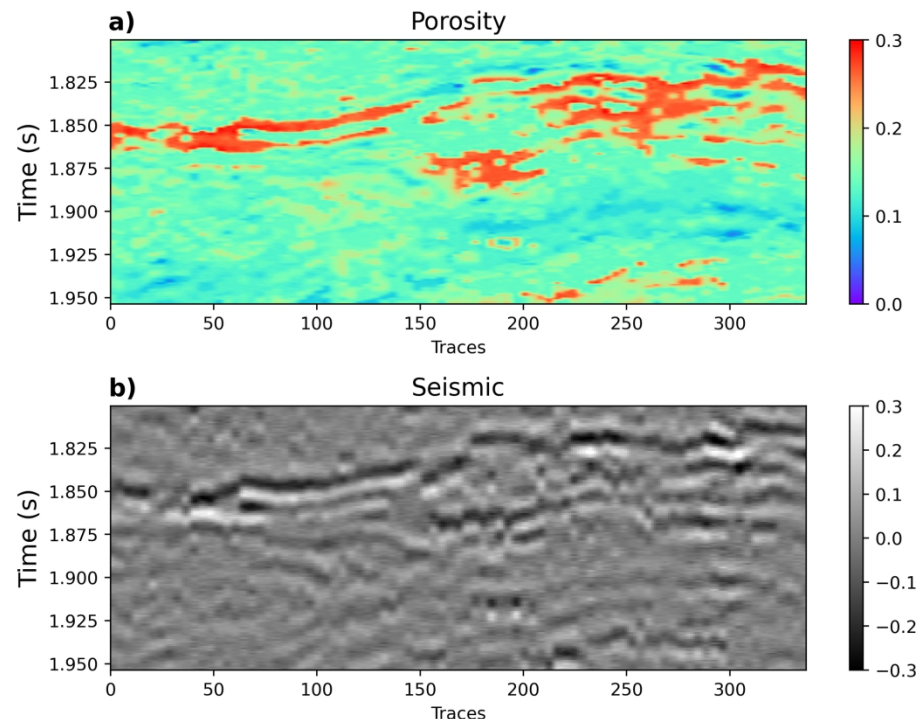


Figure 4: 2D dataset for the univariate inverse problem, a) porosity section, and b) synthetic seismic section.

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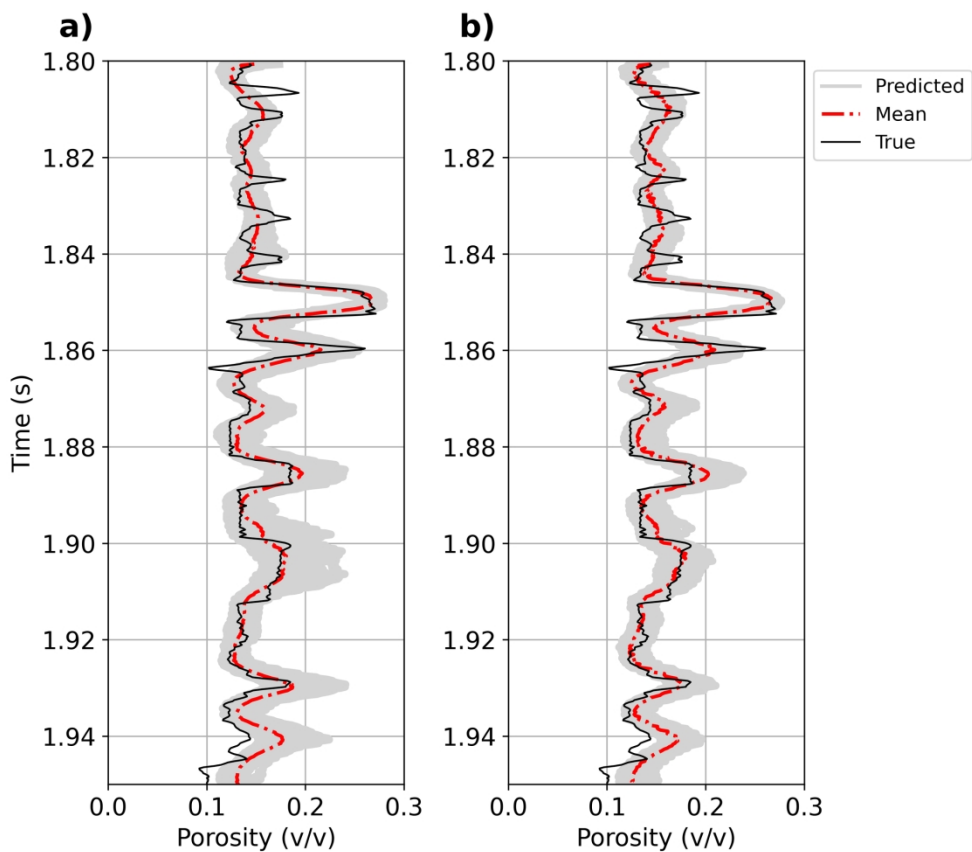


Figure 5: 1D Inversion results for the univariate problem, a) predictions from regular NN, and b) predictions from P-PINN.

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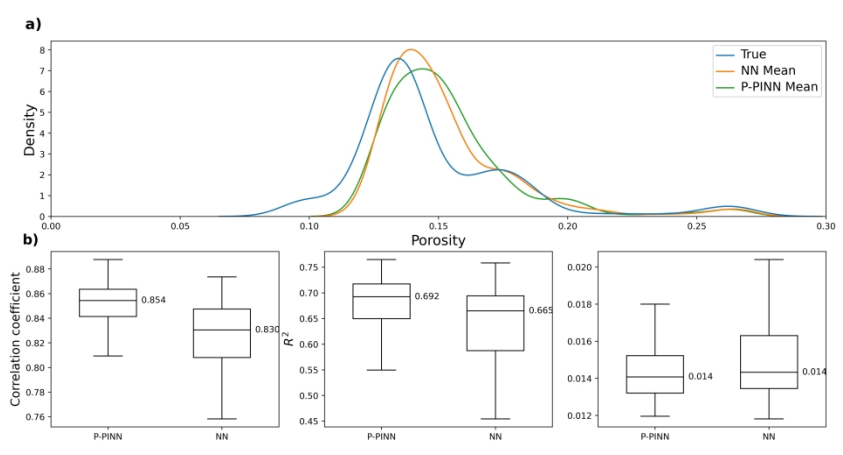


Figure 6: Comparison of metrics distributions (correlation coefficient,  $R^2$ , and MAE) between true and estimated models for regular NN and from P-PINN of the univariate inverse problem for one trace.

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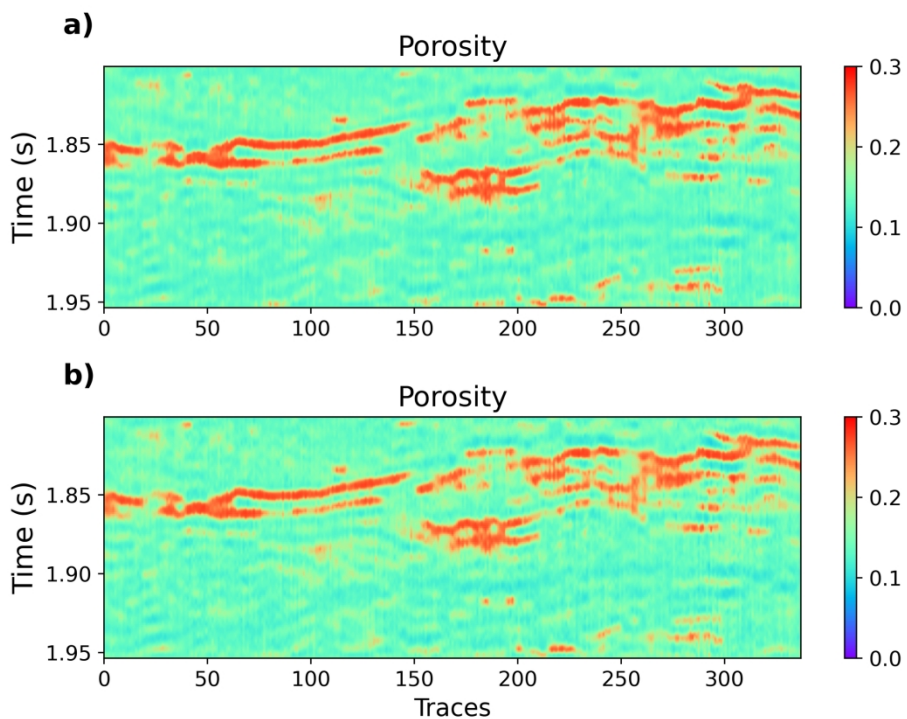


Figure 7: Mean of 2D accepted inversion results of the univariate inverse problem setting, a) predictions from regular NN, and b) predictions from P-PINN.

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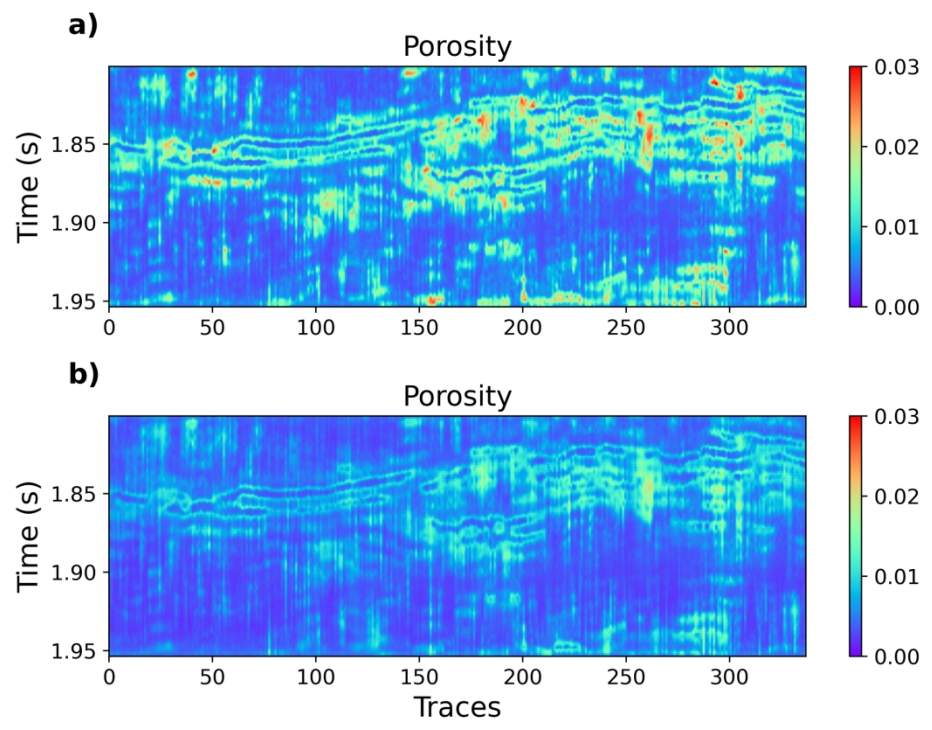


Figure 8: Standard deviation of the accepted inversion results of the univariate inverse problem, a) predictions from regular NN, and b) predictions from P-PINN.

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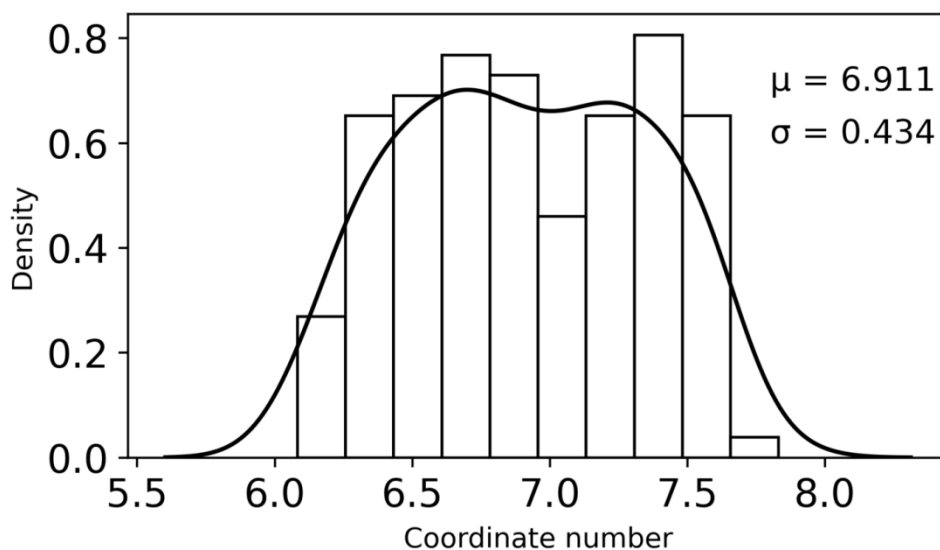


Figure 9: Distribution of estimated rock physics parameter (coordination number) of the univariate inverse problem.

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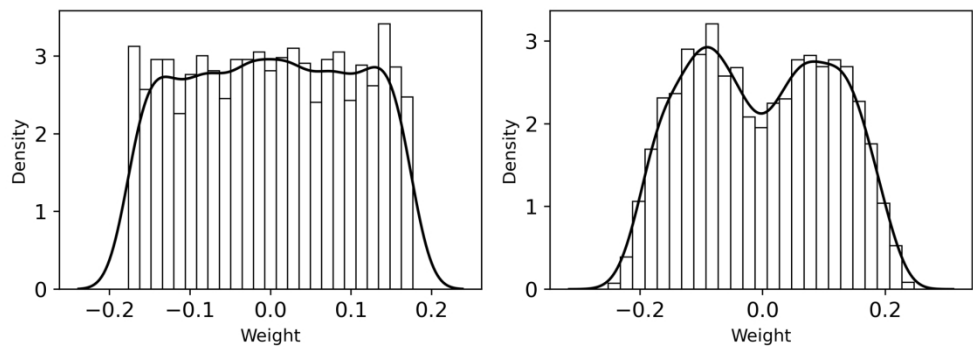


Figure 10: Posterior distributions of weights from the last regression layer: a) weights from regular NN and b) weights from P-PINN

203x76mm (300 x 300 DPI)

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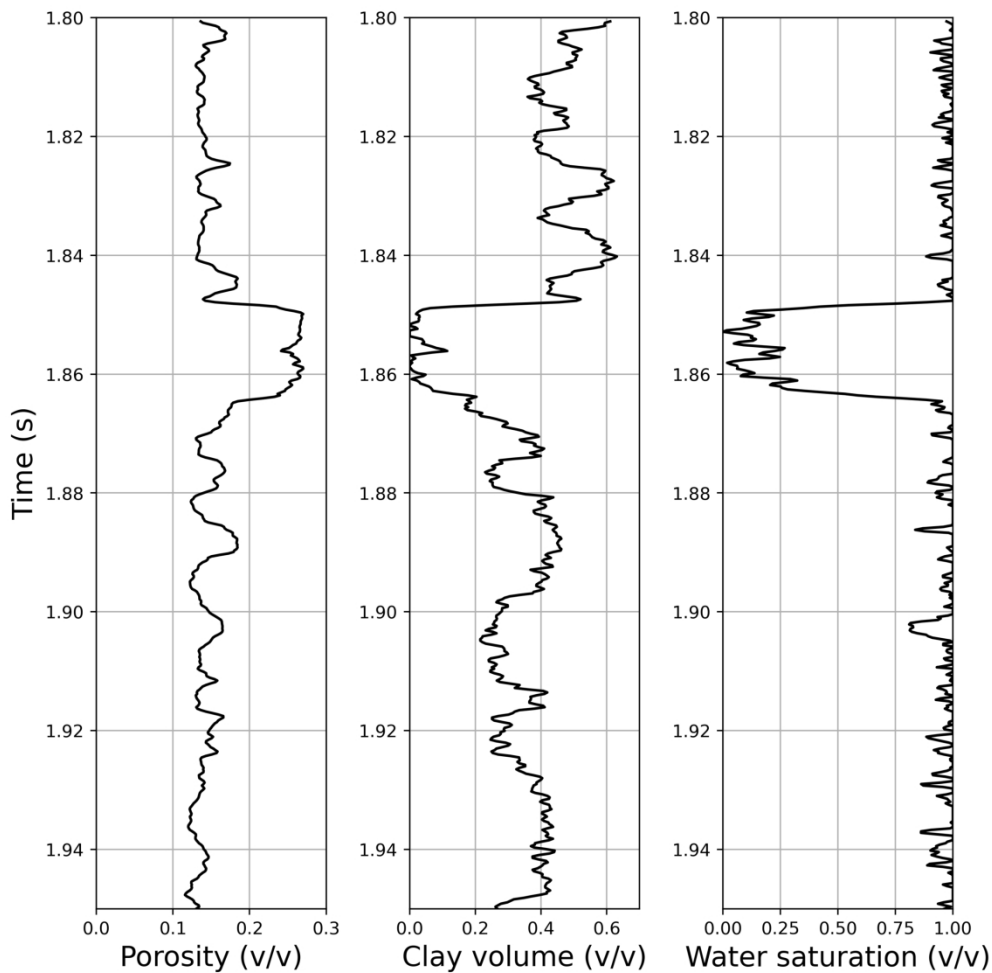


Figure 11: 1D model variables for the multivariate inverse problem, from left to right: porosity, clay volume, and water saturation.

203x203mm (300 x 300 DPI)

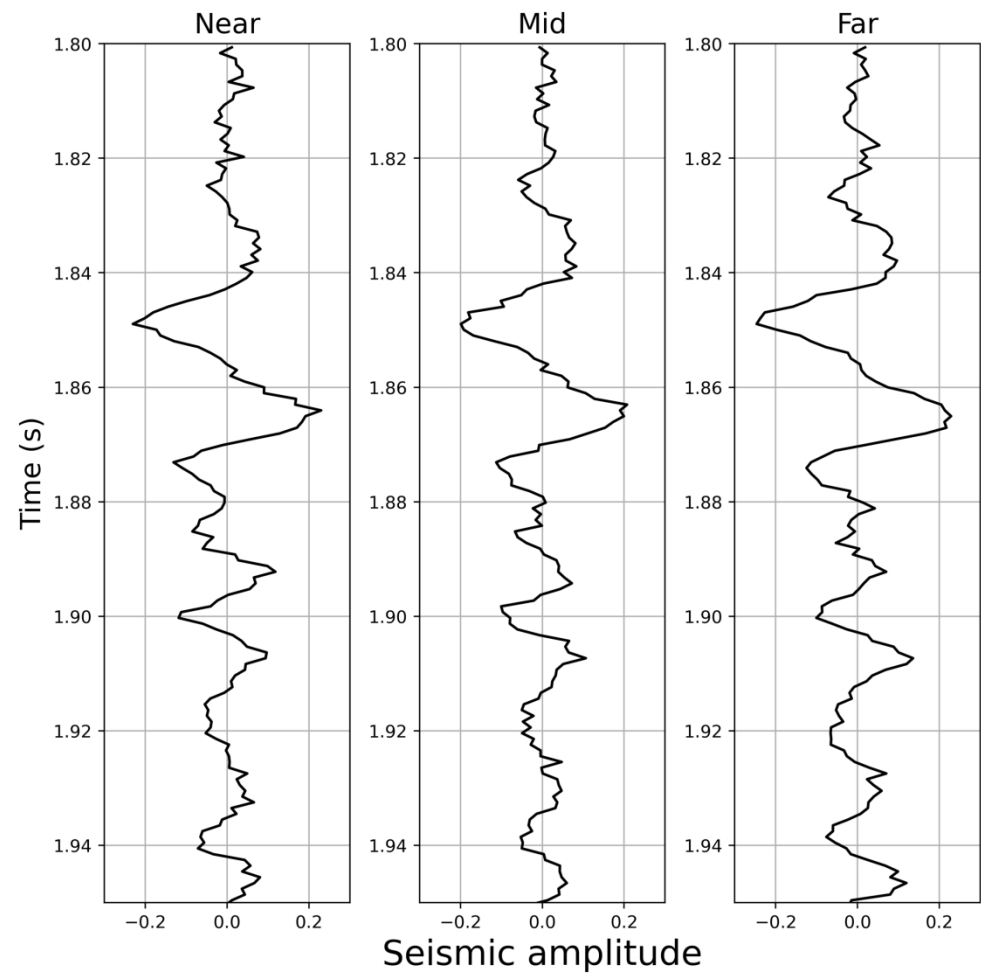


Figure 12: 1D synthetic seismic data with three partial angle stacks of the multivariate inverse problem, from left to right: near (15°), mid (30°) and far (45°).

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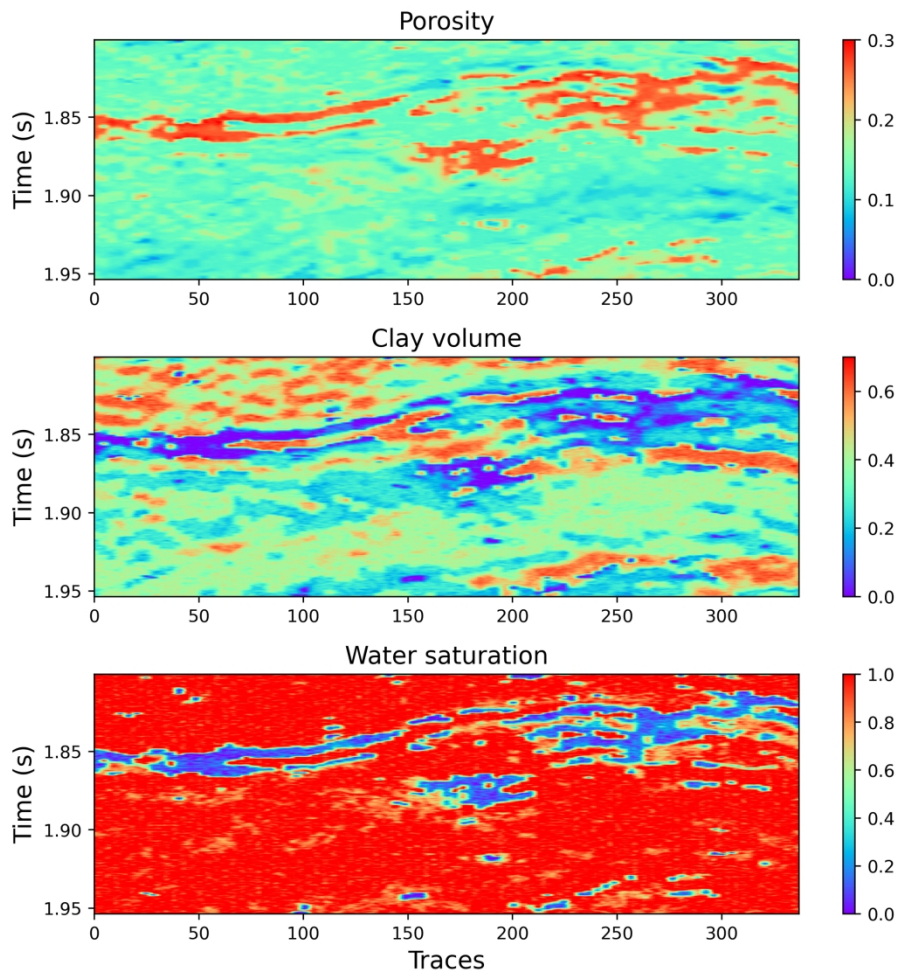


Figure 13: 2D true petrophysical variables of the multivariate inverse problem, from top to bottom: porosity, clay volume, and water saturation.

203x203mm (300 x 300 DPI)

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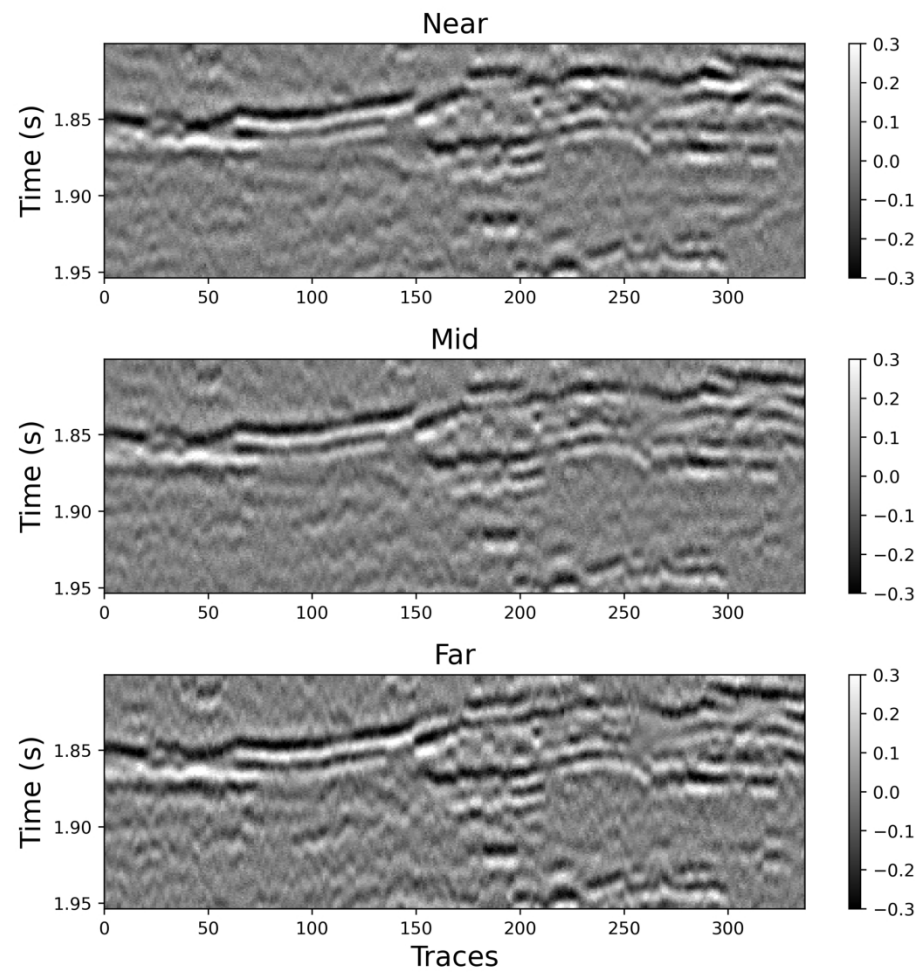


Figure 14: 2D synthetic seismic sections with three partial angles of the multivariate inverse problem, from top to bottom: near (15°), mid (30°) and far (45°).

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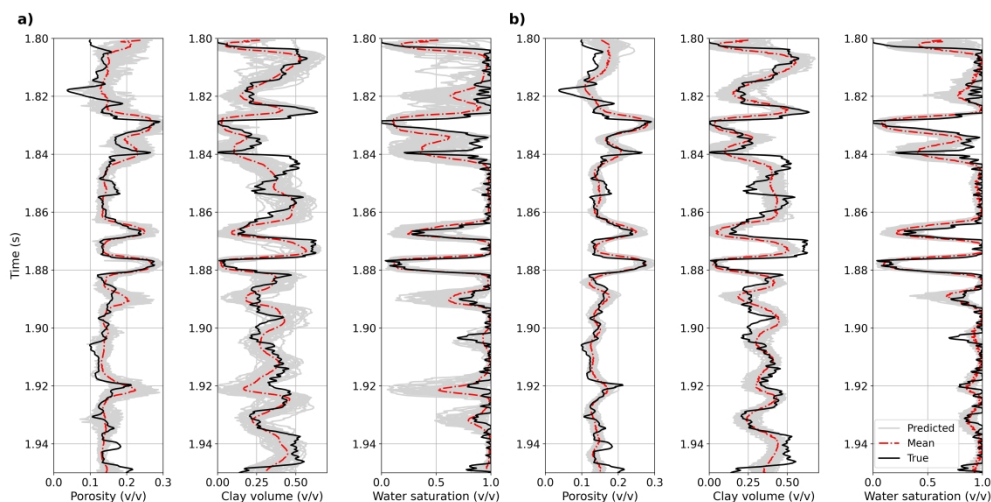


Figure 15: 1D Inversion results of the multivariate inverse problem, a) predictions from regular NN, and b) predictions from P-PINN.

487x244mm (300 x 300 DPI)

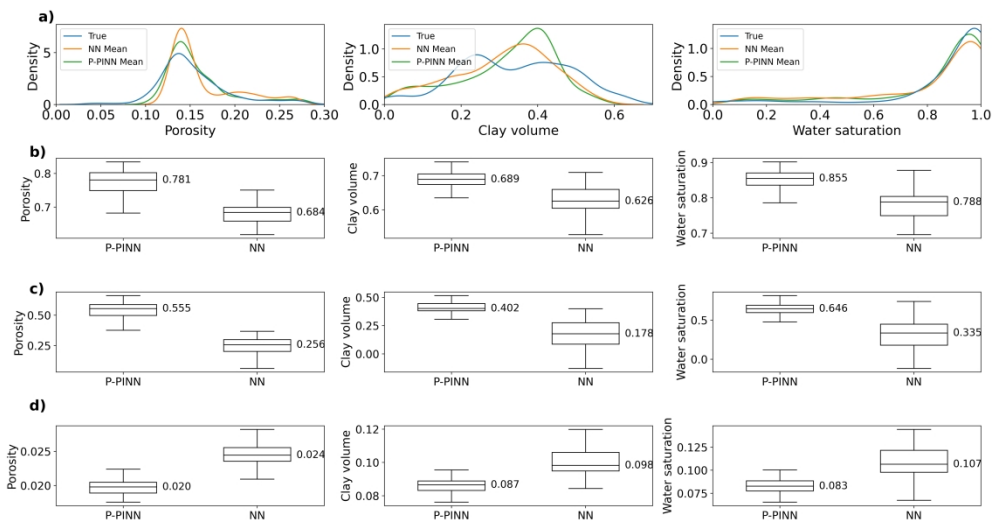


Figure 16: Comparison of metrics distributions (from top to bottom: a) Correlation coefficient, b)  $R^2$ , and c) MAE) between true and estimated models for regular NN and from P-PINN of the multivariate inverse problem for one trace.

488x257mm (300 x 300 DPI)

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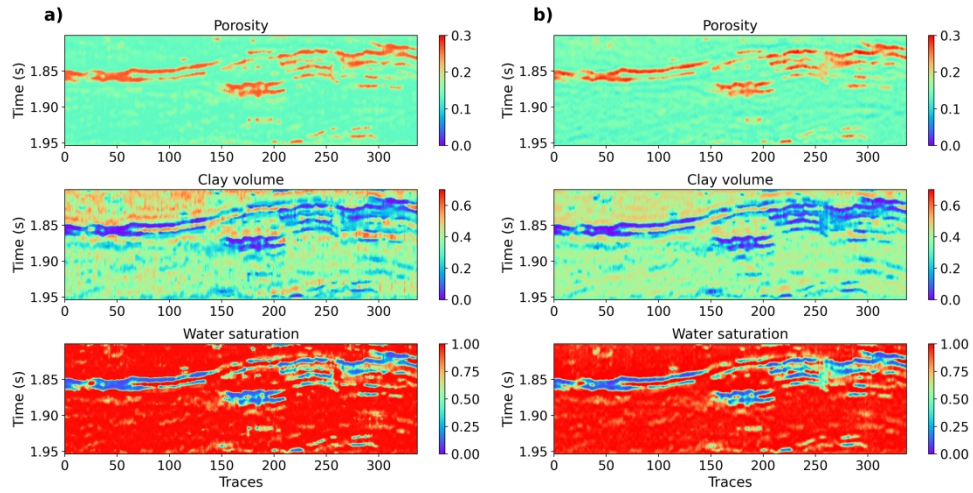


Figure 17: Mean of 2D accepted inversion results of the multivariate inverse problem, a) predictions from regular NN, and b) predictions from P-PINN.

406x203mm (300 x 300 DPI)

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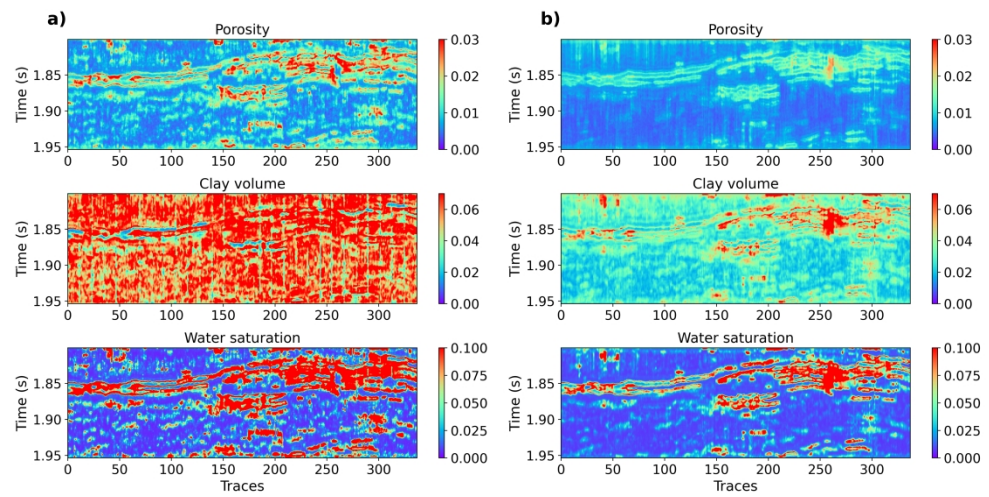


Figure 18: Standard deviation of the accepted inversion results of the multivariate inverse problem, a) predictions from regular NN, and b) predictions from P-PINN.

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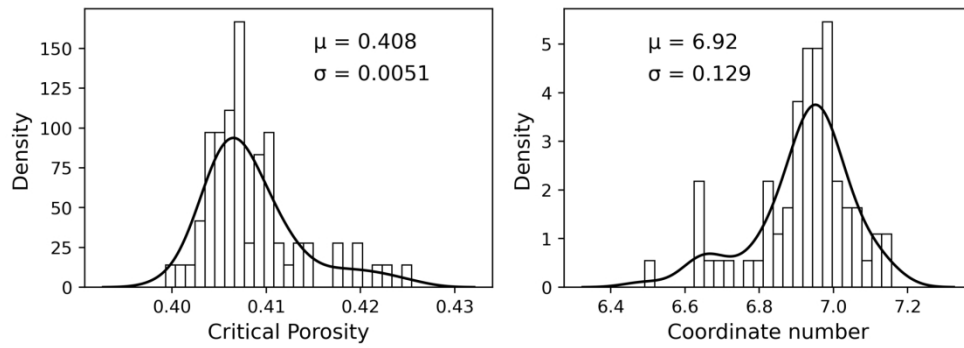


Figure 19: Distributions of estimated unknown rock physics hyperparameters of the multivariate inverse problem: a) estimated critical porosity, and b) estimated coordination number.

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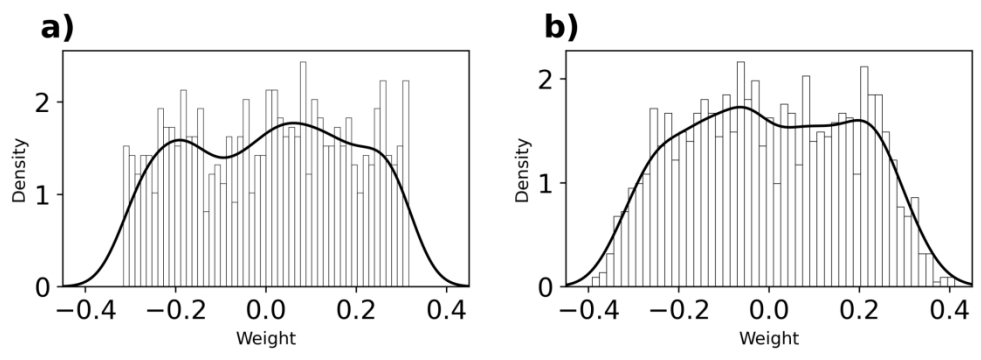


Figure 20: Posterior distributions of weights from the last regression layer: a) weights from regular NN and b) weights from P-PINN.

203x76mm (300 x 300 DPI)



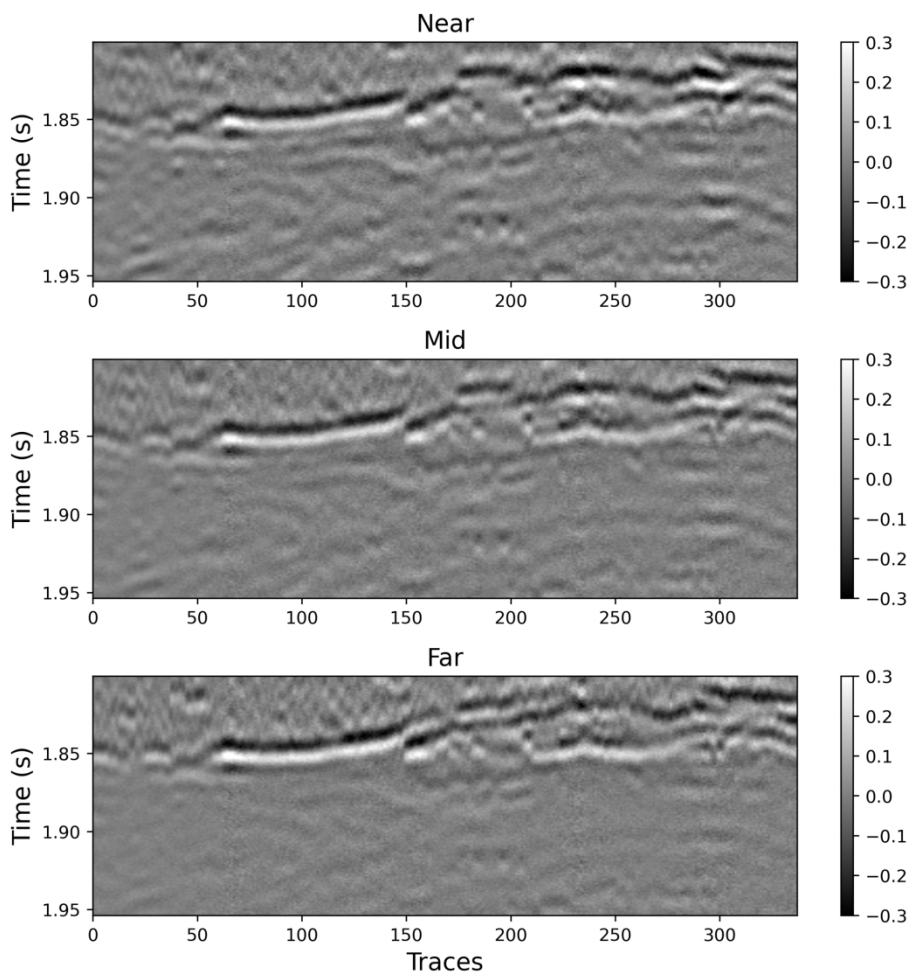


Figure 21: Synthetic seismic data with oil-water contact (OWC) at 1.85s.

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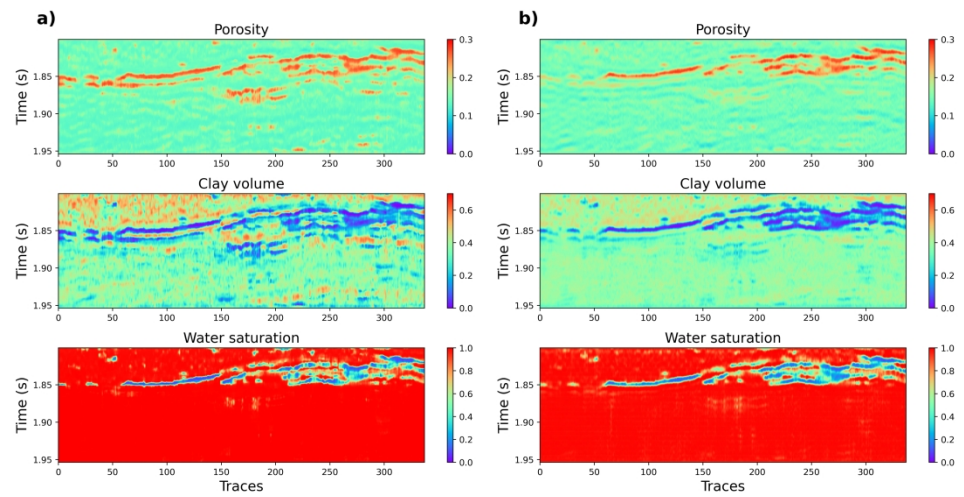


Figure 22: Inversion results for the synthetic case with OWC at 1.85s: a) means from regular NN, and b) means from P-PINN.

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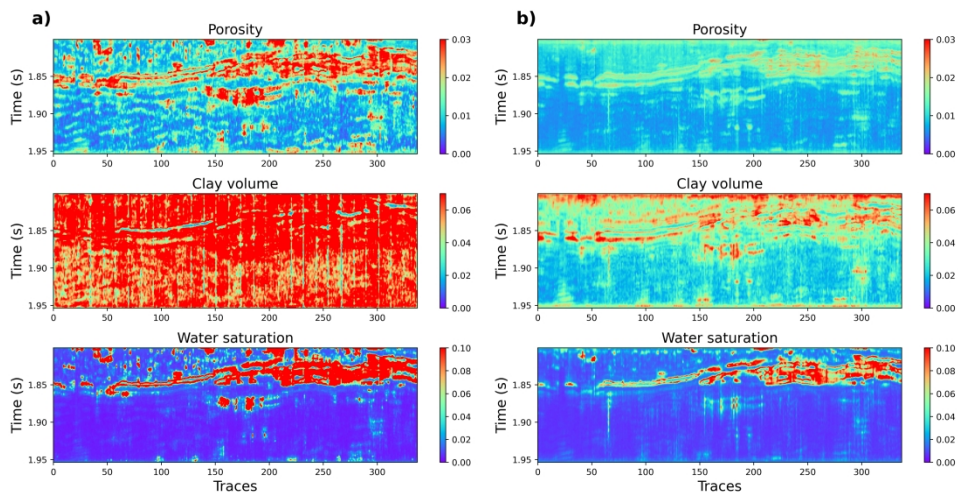


Figure 23: Uncertainty of inversion results for the synthetic case with OWC at 1.85s: a) standard deviations from regular NN, and b) standard deviations from P-PINN.

406x203mm (300 x 300 DPI)

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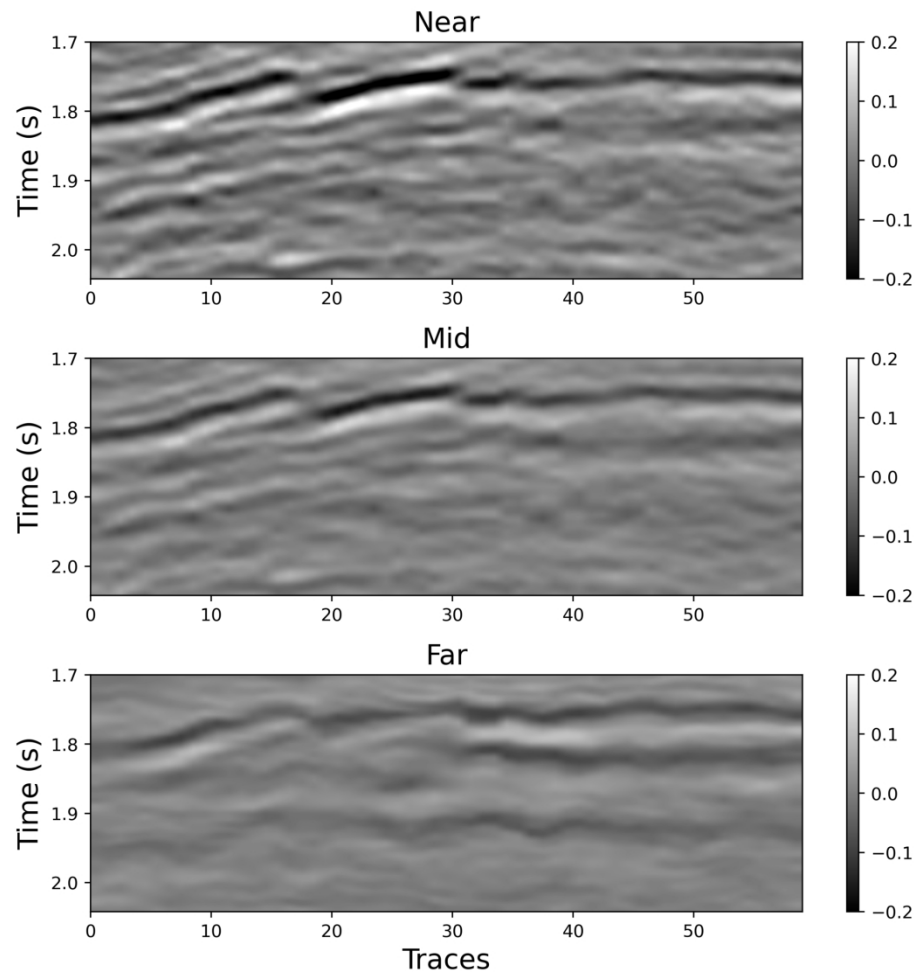


Figure 24: Seismic data for the real case application. The dataset includes three partial angle stacks: Near (15°), Mid (30°) and Far (45°)  
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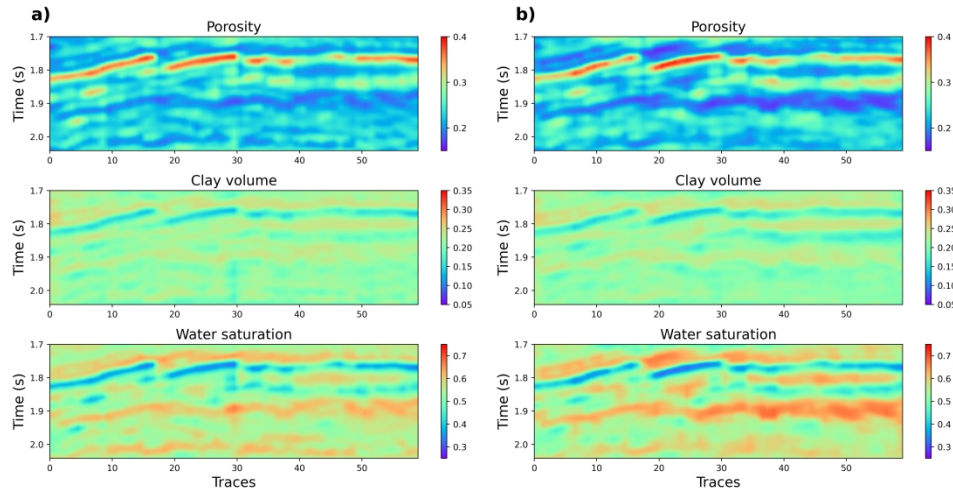


Figure 25: Inversion results for the real case application: a) means from regular NN, and b) means from P-PINN.

406x203mm (300 x 300 DPI)

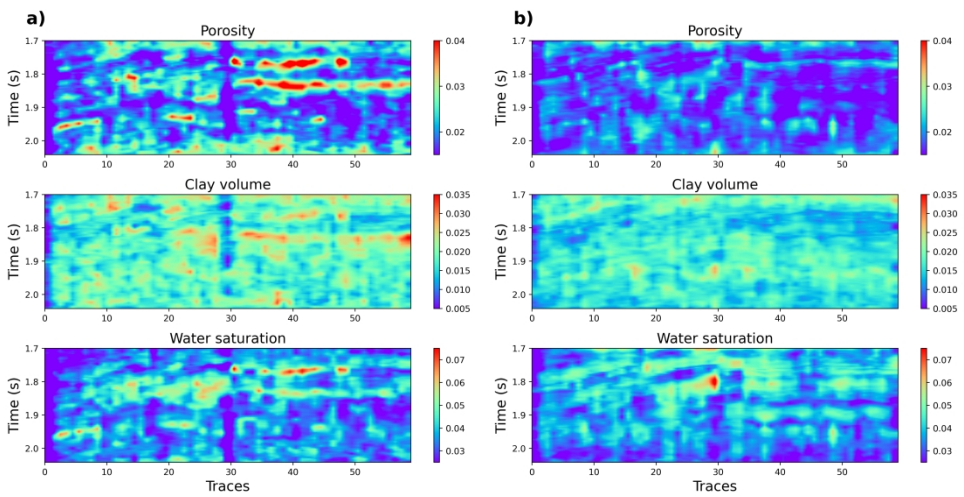


Figure 26: Uncertainty of inversion results for the real case application: a) standard deviations from regular NN, and b) standard deviations from P-PINN.

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## DATA AND MATERIALS AVAILABILITY

Data associated with this research are available and can be obtained by contacting the corresponding author.

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DOI: 10.1190/geo2023-0214.1